



Development of aroCACM/MPMPO 1.0: a model to simulate secondary organic aerosol from aromatic precursors in regional models

Matthew L. Dawson¹, Jialu Xu^{2,3}, Robert J. Griffin², and Donald Dabdub¹

¹Department of Mechanical and Aerospace Engineering, University of California, Irvine, CA 92697, USA

²Department of Civil and Environmental Engineering, Rice University, Houston, TX 77005, USA

³ACS Engineering, 16225 Park Ten Place Suite 110, Houston, TX, 77084, USA

Correspondence to: Donald Dabdub (ddabdub@uci.edu)

Received: 3 October 2015 – Published in Geosci. Model Dev. Discuss.: 15 January 2016

Revised: 27 April 2016 – Accepted: 10 May 2016 – Published: 10 June 2016

Abstract. The atmospheric oxidation of aromatic compounds is an important source of secondary organic aerosol (SOA) in urban areas. The oxidation of aromatics depends strongly on the levels of nitrogen oxides (NO_x). However, details of the mechanisms by which oxidation occurs have only recently been elucidated. Xu et al. (2015) developed an updated version of the gas-phase Caltech Atmospheric Chemistry Mechanism (CACM) designed to simulate toluene and *m*-xylene oxidation in chamber experiments over a range of NO_x conditions. The output from such a mechanism can be used in thermodynamic predictions of gas–particle partitioning leading to SOA. The current work reports the development of a model for SOA formation that combines the gas-phase mechanism of Xu et al. (2015) with an updated lumped SOA-partitioning scheme (Model to Predict the Multi-phase Partitioning of Organics, MPMPO) that allows partitioning to multiple aerosol phases and that is designed for use in larger-scale three-dimensional models. The resulting model is termed aroCACM/MPMPO 1.0. The model is integrated into the University of California, Irvine – California Institute of Technology (UCI-CIT) Airshed Model, which simulates the South Coast Air Basin (SoCAB) of California. Simulations using 2012 emissions indicate that “low- NO_x ” pathways to SOA formation from aromatic oxidation play an important role, even in regions that typically exhibit high- NO_x concentrations.

1 Introduction

Atmospheric aerosol particles negatively affect human health, contribute to reduced visibility and impact Earth’s climate through their ability to scatter and absorb radiation and affect cloud properties (Finlayson-Pitts and Pitts, 2000; Seinfeld and Pandis, 2006). Secondary organic aerosol (SOA) accounts for a significant fraction of atmospheric aerosol mass (Hallquist et al., 2009; Kanakidou et al., 2005; Kroll and Seinfeld, 2008). In urban areas, the gas-phase oxidation of aromatic hydrocarbons is thought to be an important contributor to SOA formation, although the fundamental mechanisms have only recently been elucidated (Im et al., 2014; Ng et al., 2007; Song et al., 2005; Xu et al., 2015).

In urban areas, aromatics are an important constituent of the mix of volatile organic compounds (VOCs) and are emitted primarily from industrial operations, fuel evaporation and vehicle exhaust, with smaller contributions from biomass burning and other sources (Karl et al., 2009). In addition, there is evidence of biogenic sources of aromatic compounds that may be important in rural areas (Gratien et al., 2011; White et al., 2009).

The atmospheric oxidation of aromatics is thought to proceed primarily by reaction with hydroxyl radical (OH), but nitrate (NO_3) and chlorine (Cl) radicals readily oxidize some substituted aromatics as well (Finlayson-Pitts and Pitts, 2000). The oxidation pathway (addition to the ring as opposed to hydrogen atom extraction from a side-chain) depends on the oxidant and the extent of substitution. Once initiated, oxidation proceeds through a variety of pathways

that depend strongly on the level of nitrogen oxides (NO_x) that are present. In most cases, SOA yields from aromatics are negatively correlated with NO_x concentrations (Ng et al., 2007; Song et al., 2005; Xu et al., 2015), though exceptions have been reported (Im et al., 2014).

In order to simulate the formation of SOA and other secondary pollutants such as ozone, air quality models require chemical mechanisms that predict gas-phase oxidation chemistry over a wide range of NO_x concentrations in a computationally efficient manner. Such mechanisms can be highly reduced (lumped) or highly specific depending on computational demands and application. For example, the computationally efficient mechanism of Carter et al. (2012) uses non-specific reactions to simulate the formation of secondary products, but does not track oxidation products specifically. On the other hand, Im et al. (2014) developed the comprehensive UNIPAR (Unified Partitioning-Aerosol phase Reaction) model to predict SOA formation from toluene and 1,2,3-trimethylbenzene. UNIPAR uses the detailed Master Chemical Mechanism (MCM; Jenkin et al., 2003; Saunders et al., 2003) to simulate aromatic SOA formation based on a lumped equilibrium-partitioning scheme that accounts for liquid–liquid phase separation, aqueous-phase oligomerization and organo-sulfate formation. The UNIPAR output in conjunction with outdoor chamber experiments performed over a range of NO_x conditions suggests an important role for aerosol-phase oligomerization in aromatic SOA formation.

As an intermediate approach to the non-specific and comprehensive methods mentioned above, Xu et al. (2015) updated the Caltech Atmospheric Chemical Mechanism (CACM) to include SOA formation from the gas-phase oxidation of toluene and *m*-xylene and evaluated the model through comparison to chamber experiments. When combined with a partitioning model based on absorption theory (Pankow, 1994), the updated CACM simulated SOA formation accurately over a range of NO_x conditions, without the need for separate low- and high- NO_x parameters. In addition, the reduced complexity of CACM compared to the fully explicit MCM decreases computational cost, making it ideal for incorporation into larger-scale models.

This work reports results from the development of an updated SOA-formation mechanism, aroCACM/MPMPO 1.0, based on the initial work of Xu et al. (2015) and an updated version of the Model to Predict the Multi-phase Partitioning of Organics (MPMPO; Griffin et al., 2003). The test bed for this mechanism is the University of California, Irvine – California Institute of Technology (UCI-CIT) Airshed Model, which simulates the South Coast Air Basin (SoCAB) of California. This work is presented in two parts, each with the aim of improving SOA predictions in regional models covering a range of NO_x conditions. First, the vapor pressure calculations for all lumped SOA species in the model were updated using the SIMPOL.1 method of Pankow and Asher (2008), and lumped SOA species were redistributed based

on thermodynamic properties and relative abundance. Second, the updated CACM of Xu et al. (2015) was incorporated along with appropriate lumped SOA species and a treatment of aqueous-phase partitioning of aromatic-derived SOA.

2 Model Description

2.1 UCI-CIT Airshed Model

The UCI-CIT model is an Eulerian chemical transport model that solves simultaneously the advection–diffusion equation coupled with emission, deposition and chemistry terms. The model domain is divided horizontally into an 80×30 grid of $5 \text{ km} \times 5 \text{ km}$ cells and vertically into five layers reaching a height of 1100 m above the surface. The domain includes the entire SoCAB, which includes Orange County and parts of Los Angeles, Ventura, San Bernardino and Riverside counties. The advection–diffusion equation is solved using the Quintic Splines Taylor Series Expansion approach described by Nguyen and Dabdub (2001). Deposition is based on the dry deposition flux described by Wesely for gas-phase species (Wesely, 1989), and a combination of diffusional deposition and gravitational settling for particles (Griffin et al., 2002a).

The current version of the UCI-CIT model utilizes the previous versions of CACM and MPMPO to describe formation of secondary pollutants, including SOA (Griffin et al., 2005, 2002b, a; Pun et al., 2002). Advanced numerical algorithms are used to solve the nonlinear system of highly coupled differential equations involved in a parallel computational architecture (Dabdub and Seinfeld, 1996, 1994; Nguyen and Dabdub, 2002, 2001).

2.2 Emissions and meteorology

Evaluation of aroCACM/MPMPO 1.0 described subsequently is performed with a 3-day simulation using meteorological data representative of the meteorology in the SoCAB (Carreras-Sospedra et al., 2010). It is important to note that in the testing and development presented here, a specific historical event is not being simulated.

As part of this work, the UCI-CIT model was updated to accept emissions based on the 2012 Air Quality Management Plan (AQMP) provided by the South Coast Air Quality Management District (SCAQMD, 2013), which include current (2008) emissions. All plots shown are for the second simulation day with 2008 emissions.

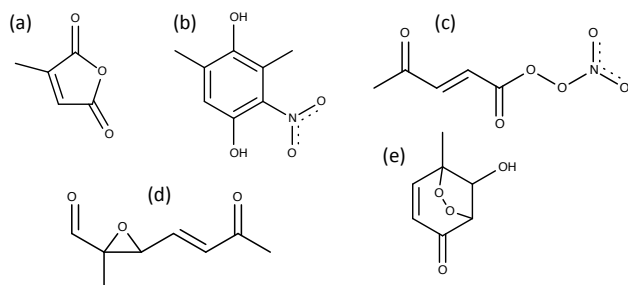
2.3 Updated vapor pressure calculations and new SOA lumping scheme

In the UCI-CIT model, the most recent version of MPMPO is used to calculate gas–particle conversion of secondary organic species (Griffin et al., 2005, 2003). In species-specific equilibrium-partitioning-based models, including MPMPO,

Table 1. New SOA groups from addition of toluene and *m*-xylene oxidation mechanisms.

SOA group	Description	Species included ^a
C1	furanones	UR24, UR72 and UR75 ^b
C2	nitro-phenols	RPR4, RP98, UR22, UR57, UR58, UR65, UR66 and UR77 ^b
C3	peroxy-nitrates	PN11 ^b , PN12, PN13 and PN14
C4	epoxides	R102 ^b and UR76
C5	peroxy-bicyclic rings	RP29, RP30 ^b and RP31

^a See Xu et al. (2015) for descriptions of individual SOA species. ^b Surrogate species.

**Figure 1.** Surrogate species for the new aromatic-derived SOA groups: (a) C1, (b) C2, (c) C3, (d) C4 and (e) C5.

thermodynamic properties including sub-cooled liquid vapor pressure (p_L^O) are required to determine the extent of gas–particle conversion (Pankow, 1994). However, experimental values for p_L^O are often unavailable, and methods for estimating p_L^O based on molecular structure and/or other properties are required. The SIMPOL.1 (simplified p_L^O prediction method.) group-contribution method of Pankow and Asher (2008) is adopted here to calculate updated SOA species vapor pressures for use in MPMPO. This method was selected because it includes contributions from functional groups that are present in the SOA species used in MPMPO but were not able to be included using the previously utilized method to calculate p_L^O (Myrdal and Yalkowsky, 1997). In addition, SIMPOL.1 was parameterized and tested using a large set of experimental p_L^O data that spans 14 orders of magnitude (Pankow and Asher, 2008).

Due to computational limitations, individual SOA species of similar molecular structure are often “lumped” into groups for the purpose of calculating gas–particle conversion in equilibrium-partitioning-based models when applied in three-dimensional air quality models. Pun et al. (2002) developed the original SOA lumping scheme used in the UCI-CIT model. Further modifications to the scheme were made by Griffin et al. (2005). Following the inclusion of SIMPOL.1 for p_L^O calculations, the existing SOA lumping scheme used by MPMPO was re-evaluated based on the structure, calculated thermodynamic parameters and relative abundance of the individual (i.e., non-lumped) SOA species. The motiva-

tion behind the specific changes made as part of this work, and their effect on model output are described in Sect. 3.1.

2.4 Aromatic-derived SOA formation

The gas-phase mechanism for toluene and *m*-xylene oxidation under varying NO_x conditions originally developed by Xu et al. (2015) was incorporated into the UCI-CIT model without modification. However, gas–particle conversion for the new aromatic-derived SOA species has been adapted for inclusion into a large-scale three-dimensional model. Because of computational limitations in such models, a SOA lumping scheme was applied to the aromatic-derived SOA species, as in the previous version of MPMPO. The new species are divided into five SOA groups labeled C1–C5. Table 1 describes the new SOA groups and lists the species and surrogates for each group. Structures for these surrogate species are shown in Fig. 1. Individual species included in the new SOA groups were selected based on structure and relative yields in the zero-dimensional model simulations of Xu et al. (2015).

As with the pre-existing SOA species, p_L^O for the new surrogate SOA species are calculated using the SIMPOL.1 method described in Sect. 2.3. These new, highly oxygenated SOA groups also could be expected to partition favorably into the aqueous phase. The MPMPO is a fully coupled mixed-phase SOA model that accounts for partitioning of each SOA group into an aqueous and organic aerosol phase simultaneously (Griffin et al., 2005, 2003). Therefore, Henry’s law constants for the new species were calculated using the method of Suzuki et al. (1992) for use in the aqueous-phase partitioning module of MPMPO. Finally, due to the new grouping scheme, the recalculated p_L^O , and the introduction of aqueous-phase partitioning for the new SOA species, the experimentally determined correction factors for p_L^O used by Xu et al. (2015) were not adopted in aroCACM/MPMPO. This could result in an underestimation of SOA formation. To further explore this potential underestimation of SOA, several simulations were performed in which correction factors were applied to bring the SIMPOL.1-calculated p_L^O into agreement with those used by Xu et al. (2015) at 298 K. Results from these simulations are discussed in Sect. 3.2.

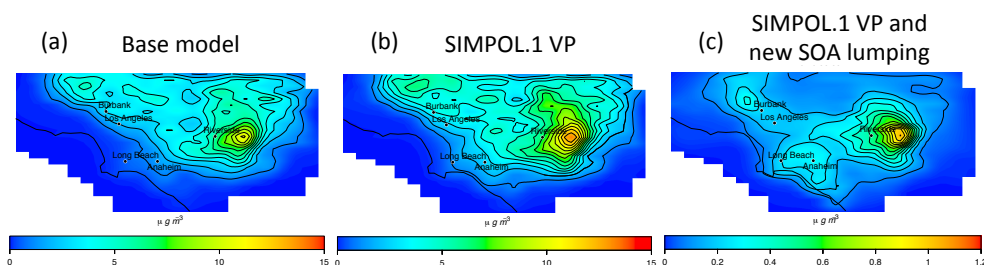


Figure 2. Modeled 24 h average total SOA concentrations (a) before updates to vapor pressure calculations and SOA lumping scheme, (b) after implementation of SIMPOL.1 for vapor pressure calculations, and (c) after adoption of SIMPOL.1 and new SOA lumping scheme. Note that the scale changes between panels.

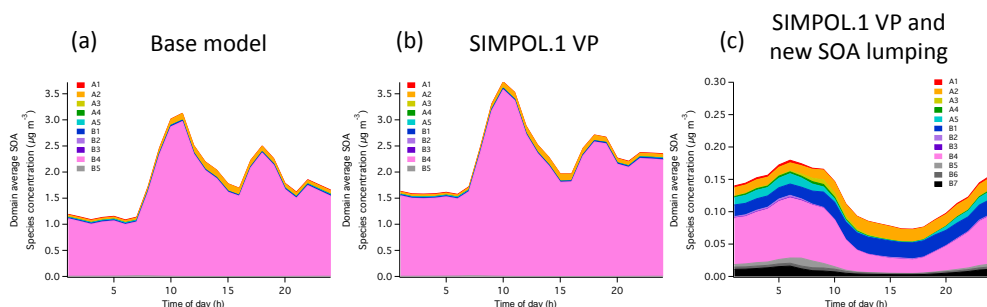


Figure 3. Modeled domain-average diurnal SOA concentrations by SOA group (a) before updates to vapor pressure calculations and SOA lumping scheme, (b) after implementation of SIMPOL.1 for vapor pressure calculations and (c) after adoption of SIMPOL.1 and new SOA lumping scheme. Note that the scale changes between panels. Surrogate structures for each of the lumped SOA species are shown in Fig. 4.

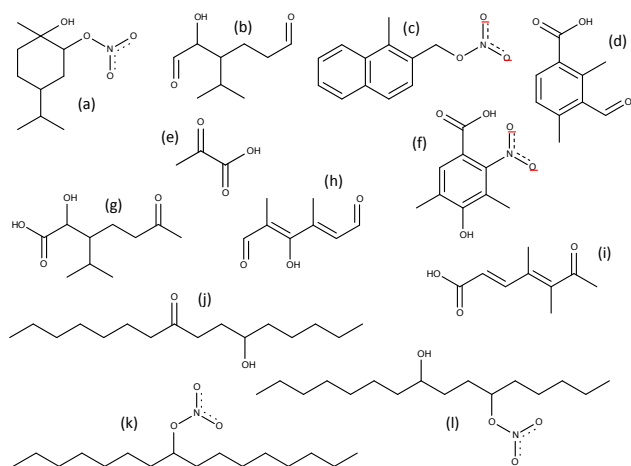


Figure 4. Surrogate species for each SOA group after changes shown in Table 2, but before addition of new aromatic-derived species: (a) B5, (b) A5, (c) B3, (d) B2, (e) A1, (f) B1, (g) A2, (h) A3, (i) A4, (j) B7, (k) B6 and (l) B4.

Emissions of aromatic compounds in the 2012 AQMP are grouped into high-SOA-yield (AROH) and low-SOA-yield (AROL) species. In the absence of more detailed speciation data, some estimation of the contribution of toluene and *m*-xylene to total aromatics must be made. To evaluate the impact of toluene and *m*-xylene oxidation on SOA concentra-

tions individually, model runs are presented in which AROH emissions are assumed to be entirely toluene or entirely *m*-xylene. A scenario where toluene and *m*-xylene are assumed to be 30 % each of AROH is also presented, as an approximation of the impact of this new chemistry in a real urban environment. Although this is a rough estimate, toluene and *m*-xylene are known to be major contributors to aromatic emissions in urban areas (de Gouw et al., 2005), and other aromatics, including *o*- and *p*-xylene, likely exhibit similar reactivity. Some aromatic compounds that are present in the atmosphere, such as styrene, are expected to have higher SOA yields due to their reaction with ozone (Hallquist et al., 2009). However, because of their relatively low concentrations, these species are not specifically treated in this work. Unless otherwise noted, all results shown for model simulations that include the new aromatic-derived SOA species use this final emissions scenario, which includes *m*-xylene and toluene.

2.5 Low-NO_x chemistry

Ashworth et al. (2015) reported several updates to CACM (what they termed CACM0.0), primarily related to low-NO_x chemistry, based on a comparison to field measurements from a forest canopy and to other gas-phase chemical mechanisms. While specific monoterpene oxidation chemistry for α -pinene, β -pinene and limonene is not included in the UCI-

Table 2. Updates to the SOA lumping scheme as part of this work, along with their motivation. The original SOA lumping scheme is described in Pun et al. (2002) and Griffin et al. (2005).

SOA group	Change	Reason
A1	made UR21 (ketopropanoic acid) the surrogate species	Species UR21 has a much higher domain average concentration than the previous surrogate, UR28 (oxalic acid).
A2, A4	moved UR3 from A4 to A2	The calculated p_L^O for UR3 is closer to species in SOA group A2.
A2, A4	moved UR26 from A2 to A4	Same as above.
A2	made UR3 the surrogate species	Species UR3 has representative p_L^O , Henry's law constant and structure for group A2 and has the highest domain average concentration.
A4	made UR26 the surrogate species	Same as above.
B4	removed UR34 from group B4	Species UR34 is no longer produced in the gas-phase mechanism (Griffin et al., 2005).
B4, B6*, B7*	split group B4 into B4 (AP12), B6 (AP11) and B7 (UR20)	These three SOA species have high domain average concentrations and calculated p_L^O (298 K) that vary by an order of magnitude or more.

* New lumped SOA species.

CIT version of CACM, two of their recommended adjustments were made. The rate constants for the reaction of organic peroxy radicals (RO_2) with hydroperoxy radicals (HO_2) or other RO_2 species were increased to be in agreement with Table A2 in Ashworth et al. (2015). Additionally, the formation of isoprene-derived organic nitrates, along with their subsequent reactions with OH, were added. However, these modifications had little effect on modeled NO and NO_2 concentrations or SOA concentration. This would be expected as biogenic VOCs play a much greater role in SOA formation in a forested region (for which their model was developed), than in the relatively high- NO_x , arid SoCAB region.

3 Results and discussion

3.1 Updated vapor pressure and new SOA lumping scheme

The change in predicted SOA upon implementation of the SIMPOL.1 method for vapor pressure calculations is shown in Figs. 2 and 3. Use of the SIMPOL.1 vapor pressure calculations results in a modest increase in SOA concentration over the base case (Fig. 2a, b). As seen in Fig. 3a and b, this increase can be attributed primarily to SOA group B4, which includes oxidation products of long-chain alkanes (Griffin et al., 2005; Pun et al., 2002). Because of the large contribution of SOA group B4 to total SOA concentration, a redistribution of SOA species was considered and performed. A summary of modifications to the SOA lumping scheme is presented in Table 2. The structures of the resulting 12 surrogate SOA species are shown in Fig. 4.

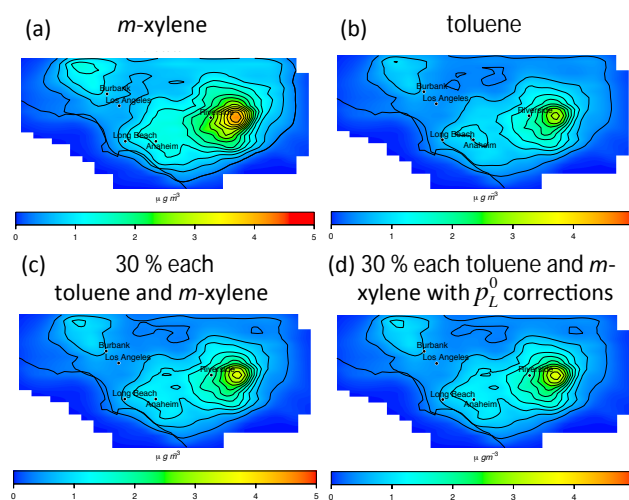


Figure 5. Modeled 24 h average SOA concentrations after inclusion of new aromatic chemistry assuming AROH emissions are (a) all *m*-xylene, (b) all toluene or (c, d) 30 % each *m*-xylene and toluene. Vapor pressures are calculated using (a, b, c) SIMPOL.1, or (d) SIMPOL.1 with SOA species C1–C5 adjusted to match those used by Xu et al. (2015) at 298 K.

The new redistribution results in a large decrease in total SOA concentration (Fig. 2c). The contribution of each SOA group to total SOA in the redistributed scheme is shown in Fig. 3c. Splitting SOA group B4 (long-chain alkane oxidation products) into groups B4, B6 and B7 (Table 2) results in an order-of-magnitude decrease in SOA from these species. Previously, the surrogate species for group B4 had both the lowest modeled vapor pressure and lowest domain-wide average concentration of all species in that SOA group. This had resulted in an over-estimation of SOA from this source,

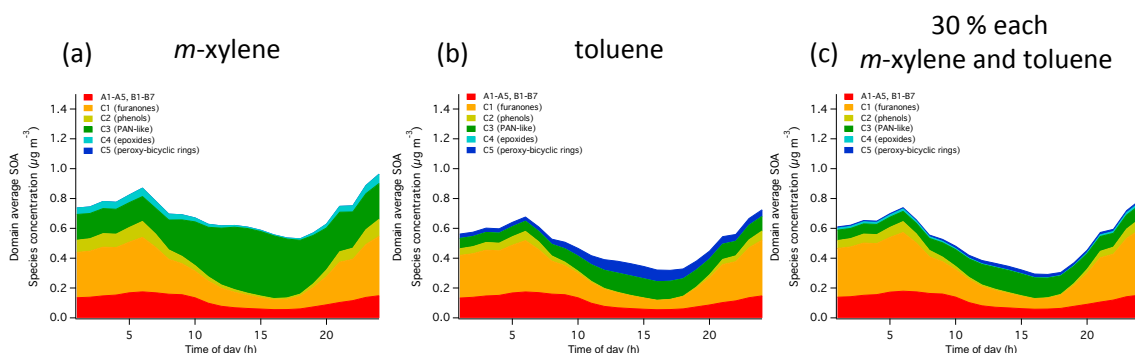


Figure 6. Modeled domain-average diurnal SOA concentrations by SOA group after inclusion of new aromatic chemistry assuming AROH emissions are (a) all *m*-xylene, (b) all toluene or (c) 30 % each *m*-xylene and toluene.

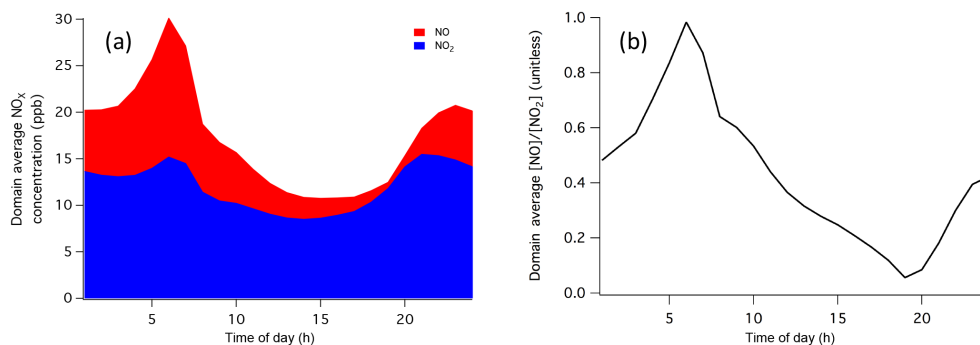


Figure 7. Modeled domain-average diurnal (a) NO and NO₂ concentrations and (b) [NO] to [NO₂] ratio, assuming AROH emissions are 30 % each *m*-xylene and toluene.

as can be seen by comparing Fig. 3b and c. The other major effect of the redistribution of SOA is the decrease in SOA formation from group A2. Although the use of species UR3 as the surrogate for SOA group A2 results in a lower modeled p_L^O for this group, the largest contribution to group A2 had come from species UR26, which has a higher p_L^O than the other species in group A2. Moving UR26 to group A4 and using it as the surrogate for this group results in a decrease in the combined SOA concentrations from groups A2 and A4.

3.2 SOA formation from toluene and *m*-xylene oxidation

Figure 5 shows 24 h average SOA concentrations assuming AROH emissions are all *m*-xylene, all toluene or 30 % each *m*-xylene and toluene (with the remaining 40 % following the original CACM AROH chemistry). Comparison with Fig. 6c shows that aromatic-derived SOA accounts for 75 % or more of total modeled SOA concentration in the domain. In these simulations, *m*-xylene oxidation leads to higher total SOA concentration than does toluene oxidation (Fig. 5a, b). The opposite trend is seen in the simulation results of Xu et al. (2015) and the experiments on which the model was based, both of which were performed under dry conditions at room temperature. Several factors may contribute to this difference

in results. First, vapor pressure correction factors used by Xu et al. (2015) were not used here. In addition, the partitioning model here allows for aqueous-phase SOA formation. Finally, it should be noted that the previous experiments and simulations allowed the parent hydrocarbon to react almost to completion. In fact, the reaction of *m*-xylene with OH occurs more rapidly than that of toluene, such that a larger fraction of *m*-xylene reacts in that simulation, potentially leading to more SOA formation despite having a smaller yield when observed in the laboratory.

As shown in Fig. 6 for the *m*-xylene and toluene scenarios, the speciated SOA compositions are in reasonable agreement with the results of Xu et al. (2015) under high-NO_x conditions. In both cases, there are strong contributions from furanones (C1) and peroxy-nitrates (C3), and in the case of toluene, peroxy-bicyclic rings (C5). However, in both the *m*-xylene and toluene cases, the relative SOA contribution from phenols (C2) is smaller than would be expected from the zero-dimensional model simulations of Xu et al. (2015) for high-NO_x conditions and corresponds more closely to low-NO_x conditions. It is important to note that these conditions are total NO_x and do not generally consider the split of NO_x between NO and NO₂.

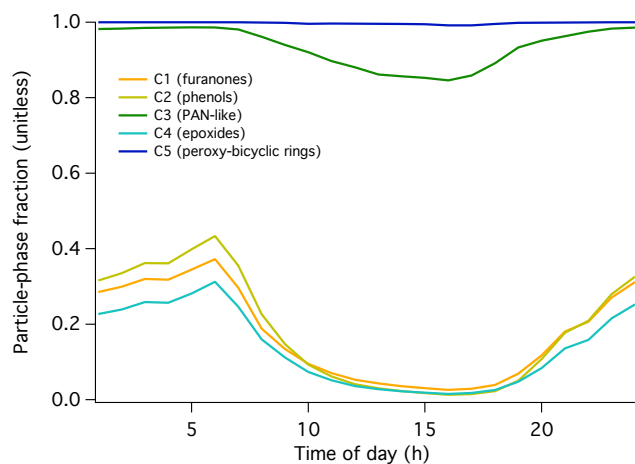


Figure 8. Diurnal fraction of aromatic-derived SOA species in the particle phase, averaged over the model domain. This fraction is defined as the mass concentration of a lumped SOA species relative to its total mass concentration (gas and particle phase).

Because the NO_x levels in the SoCAB are relatively high regardless of the time of day, modeled domain average NO and NO_2 concentrations are shown in Fig. 7a. In general, the contribution to total SOA from groups C1, C2 and C4 are correlated with high NO concentrations, and groups C3 and C5 with lower relative NO concentrations (Fig. 7b) and peak OH concentrations (early afternoon; not shown). In the gas-phase aromatic oxidation mechanism, NO competes with NO_2 to react with acyl peroxy radical species. When the $[\text{NO}]/[\text{NO}_2]$ ratio is low, the NO_2 reaction forms peroxy-nitrates and leads to an increase in group C3. When the $[\text{NO}]/[\text{NO}_2]$ ratio is high, the NO reaction with certain peroxy radical species can form furanones and lead to an increase in group C1. The reaction of NO_2 with phenolic radical species to form nitro-phenols leads to an increase in SOA formation from group C2 when NO_2 concentrations are highest. Meanwhile, NO competes with the intra-molecular cyclization reaction of certain peroxy radical species, leading to a decrease in peroxy-bicyclic ring species (SOA group C5) under high- NO conditions. The diurnal trend of SOA derived from group C4 is more difficult to explain based on the gas-phase mechanism. Some contribution to the decrease seen in groups C1, C2 and C4 may be due to the increased preference for the gas phase of these semi-volatile species during the daytime, as is supported by the average fraction of species C1, C2 and C4 in the particle phase as a function of the time of day (Fig. 8).

Although the experimental validation of the mechanism used in the current work was for dry conditions, aqueous-phase partitioning of aromatic-derived SOA species was incorporated into aroCACM/MPMPO 1.0. To evaluate the importance of aqueous-phase partitioning for the aromatic-derived SOA species in the updated UCI-CIT model, a simulation was performed in which species C1–C5 were al-

lowed to partition only to the organic phase. Results indicate that $> 99\%$ of modeled SOA from species C1–C5 is in the aqueous-phase. It is possible that the vapor pressure correction factors used by Xu et al. (2015) would increase the fraction of species C1–C5 in the organic phase. Therefore, a final simulation was performed in which the SIMPOL.1-calculated vapor pressures were adjusted to match those used by Xu et al. (2015) at 298 K. (Note that this model run includes aqueous-phase partitioning of species C1–C5.) This results in little change to modeled SOA, as seen by comparing Fig. 5c and d, again suggesting that aqueous-phase partitioning may play a dominant role in SOA formation from the aromatic-derived SOA species. (Note that in the final version of the model, the unadjusted SIMPOL.1 vapor pressures are used.)

In addition to aqueous-phase partitioning, results from other experimental and modeling work suggest that aqueous-phase oligomerization may be an important route to SOA formation from aromatics under atmospherically relevant conditions (Im et al., 2014). Future studies using experiments on aromatic oxidation under a variety of relative humidity conditions will be required to further assess the relative importance of aqueous- vs. organic-phase partitioning and the role of oligomerization and to refine the aroCACM/MPMPO 1.0 model.

4 Conclusions

This work reports the development of aroCACM/MPMPO, a gas-phase and SOA model designed for use in large-scale chemical transport models that includes important routes to SOA formation from the oxidation of aromatic species. aroCACM/MPMPO makes use of updated schemes for vapor pressure estimation and grouping SOA species and includes both aqueous- and organic-phase partitioning for all SOA species using an equilibrium-partitioning module. Results presented here from the implementation of aroCACM/MPMPO 1.0 into the UCI-CIT Airshed Model support the importance of aromatic-derived SOA in urban areas. Of particular note is the manner in which the concentrations of individual SOA-partitioning species demonstrate different trends with the time of day. Furanones, phenols, and epoxides are found to increase their contribution to total SOA during times of high NO_x and low OH , while peroxy-nitrate and peroxy-bicyclic ring species show the opposite trend. This indicates that even in areas generally considered “high” NO_x it is important to include the contribution of low- NO_x pathways to SOA formation from aromatic oxidation, primarily due to the importance of the NO to NO_x ratio.

Code availability

Model code is available by request via email request to ddabub@uci.edu.

The Supplement related to this article is available online at doi:10.5194/gmd-9-2143-2016-supplement.

Acknowledgements. This work was performed under grants from the National Science Foundation (CHE-1213623 and ATM-0901580). Any opinions, findings and conclusions or recommendations expressed in this material are those of the authors and do not necessarily reflect the views of the National Science Foundation.

Edited by: A. Lauer

References

- Ashworth, K., Chung, S. H., Griffin, R. J., Chen, J., Forkel, R., Bryan, A. M., and Steiner, A. L.: FORest Canopy Atmosphere Transfer (FORCAsT) 1.0: a 1-D model of biosphere–atmosphere chemical exchange, *Geosci. Model Dev.*, 8, 3765–3784, doi:10.5194/gmd-8-3765-2015, 2015.
- Carreras-Sospedra, M., Vutukuru, S., Brouwer, J., and Dabdub, D.: Central power generation versus distributed generation – An air quality assessment in the South Coast Air Basin of California, *Atmos. Environ.*, 44, 3215–3223, 2010.
- Carter, W., Heo, G., Cocker, D., and Nakao, S.: SOA formation: Chamber study and model development. Final report to the California Air Resources Board, Contract No. 08-326, Tech. rep., 2012.
- Dabdub, D. and Seinfeld, J. H.: Air quality modeling on massively parallel computers, *Atmos. Environ.*, 28, 1679–1687, 1994.
- Dabdub, D. and Seinfeld, J. H.: Parallel computation in atmospheric chemical modeling, *Parallel Comput.*, 22, 111–130, 1996.
- de Gouw, J. A., Middlebrook, A. M., Warneke, C., Goldan, P. D., Kuster, W. C., Roberts, J. M., Fehsenfeld, F. C., Worsnop, D. R., Canagaratna, M. R., Pszenny, A. a. P., Keene, W. C., Marchewka, M., Bertman, S. B., and Bates, T. S.: Budget of organic carbon in a polluted atmosphere: Results from the New England Air Quality Study in 2002, *J. Geophys. Res.*, 110, D16305, doi:10.1029/2004JD005623, 2005.
- Finlayson-Pitts, B. and Pitts, J.: Chemistry of the upper and lower atmosphere – Theory, experiments and applications, Academic Press, San Diego, 2000.
- Gratien, A., Johnson, S. N., Ezell, M. J., Dawson, M. L., Bennett, R., and Finlayson-Pitts, B. J.: Surprising formation of *p*-cymene in the oxidation of α -pinene in air by the atmospheric oxidants OH, O₃, and NO₃, *Environ. Sci. Technol.*, 45, 2755–2760, 2011.
- Griffin, R. J., Dabdub, D., Kleeman, M. J., Fraser, M. P., Cass, G. R., and Seinfeld, J. H.: Secondary organic aerosol 3. Urban/regional scale model of size- and composition-resolved aerosols, *J. Geophys. Res.-Atmos.*, 107, 4334, doi:10.1029/2001JD000544, 2002a.
- Griffin, R. J., Dabdub, D., and Seinfeld, J. H.: Secondary organic aerosol 1. Atmospheric chemical mechanism for production of molecular constituents, *J. Geophys. Res.*, 107, 4332, doi:10.1029/2001JD000541, 2002b.
- Griffin, R. J., Nguyen, K., Dabdub, D., and Seinfeld, J. H.: A coupled hydrophobic-hydrophilic model for predicting secondary organic aerosol formation, *J. Atmos. Chem.*, 44, 171–190, 2003.
- Griffin, R. J., Dabdub, D., and Seinfeld, J. H.: Development and initial evaluation of a dynamic species-resolved model for gas phase chemistry and size-resolved gas/particle partitioning associated with secondary organic aerosol formation, *J. Geophys. Res.*, 110, D05304, doi:10.1029/2004JD005219, 2005.
- Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, Th. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155–5236, doi:10.5194/acp-9-5155-2009, 2009.
- Im, Y., Jang, M., and Beardsley, R. L.: Simulation of aromatic SOA formation using the lumping model integrated with explicit gas-phase kinetic mechanisms and aerosol-phase reactions, *Atmos. Chem. Phys.*, 14, 4013–4027, doi:10.5194/acp-14-4013-2014, 2014.
- Jenkin, M. E., Saunders, S. M., Wagner, V., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile organic compounds, *Atmos. Chem. Phys.*, 3, 181–193, doi:10.5194/acp-3-181-2003, 2003.
- Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate modelling: a review, *Atmos. Chem. Phys.*, 5, 1053–1123, doi:10.5194/acp-5-1053-2005, 2005.
- Karl, T., Apel, E., Hodzic, A., Riemer, D. D., Blake, D. R., and Wiedinmyer, C.: Emissions of volatile organic compounds inferred from airborne flux measurements over a megacity, *Atmos. Chem. Phys.*, 9, 271–285, doi:10.5194/acp-9-271-2009, 2009.
- Kroll, J. H. and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and evolution of low-volatility organics in the atmosphere, *Atmos. Environ.*, 42, 3593–3624, 2008.
- Myrdal, P. B. and Yalkowsky, S. H.: Estimating pure component vapor pressures of complex organic molecules, *Ind. Eng. Chem. Res.*, 36, 2494–2499, 1997.
- Ng, N. L., Kroll, J. H., Chan, A. W. H., Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from *m*-xylene, toluene, and benzene, *Atmos. Chem. Phys.*, 7, 3909–3922, doi:10.5194/acp-7-3909-2007, 2007.
- Nguyen, K. and Dabdub, D.: Two-level time-marching scheme using splines for solving the advection equation, *Atmos. Environ.*, 35, 1627–1637, 2001.
- Nguyen, K. and Dabdub, D.: Semi-Lagrangian flux scheme for the solution of the aerosol condensation/evaporation equation, *Aerosol Sci. Tech.*, 36, 407–418, 2002.
- Pankow, J. F.: An absorption model of the gas/aerosol partitioning involved in the formation of secondary organic aerosol, *Atmos. Environ.*, 28, 189–193, 1994.

- Pankow, J. F. and Asher, W. E.: SIMPOL.1: a simple group contribution method for predicting vapor pressures and enthalpies of vaporization of multifunctional organic compounds, *Atmos. Chem. Phys.*, 8, 2773–2796, doi:10.5194/acp-8-2773-2008, 2008.
- Pun, B. K., Griffin, R. J., Seigneur, C., and Seinfeld, J. H.: Secondary organic aerosol 2. Thermodynamic model for gas/particle partitioning of molecular constituents, *J. Geophys. Res.*, 107, 4333, doi:10.1029/2001JD000542, 2002.
- Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic volatile organic compounds, *Atmos. Chem. Phys.*, 3, 161–180, doi:10.5194/acp-3-161-2003, 2003.
- SCAQMD: Final 2012 Air Quality Management Plan, Tech. rep., South Coast Air Quality Management District, Diamond Bar, California, 2013.
- Seinfeld, J. and Pandis, S.: *Atmospheric chemistry and physics: From air pollution to climate change*, Wiley-Interscience, Hoboken, New Jersey, 2006.
- Song, C., Na, K., and Cocker, D. R.: Impact of the hydrocarbon to NO_x ratio on secondary organic aerosol formation, *Environ. Sci. Technol.*, 39, 3143–3149, 2005.
- Suzuki, T., Ohtaguchi, K., and Koide, K.: Application of principal components analysis to calculate Henry's constant from molecular structure, *Comput. Chem.*, 16, 41–52, 1992.
- Wesely, M.: Parameterization of surface resistances to gaseous dry deposition in regional-scale numerical models, *Atmos. Environ.*, 23, 1293–1304, 1989.
- White, M. L., Russo, R. S., Zhou, Y., Ambrose, J. L., Haase, K., Frinak, E. K., Varner, R. K., Wingenter, O. W., Mao, H., Talbot, R., and Sive, B. C.: Are biogenic emissions a significant source of summertime atmospheric toluene in the rural Northeastern United States?, *Atmos. Chem. Phys.*, 9, 81–92, doi:10.5194/acp-9-81-2009, 2009.
- Xu, J., Griffin, R. J., Liu, Y., Nakao, S., and Cocker III, D. R.: Simulated impact of NO_x on SOA formation from oxidation of toluene and *m*-xylene, *Atmos. Environ.*, 101, 217–225, 2015.